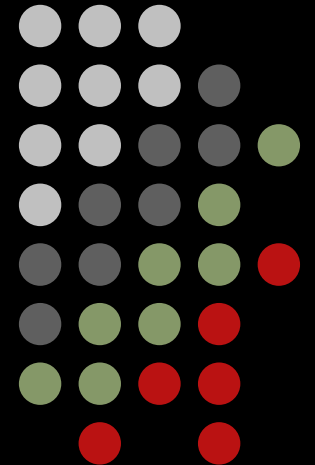


Underground reduction of Gravity Gradient Noise.



Introduction



In future advanced interferometers Newtonian (Gravity Gradient) Noise will be one of the fundamental limitations for the sensitivity in the low frequency region.

- × Can it be estimated?
- × What are the most important sources?
- × Can it be reduced?
- × What we gain going underground?

Outline



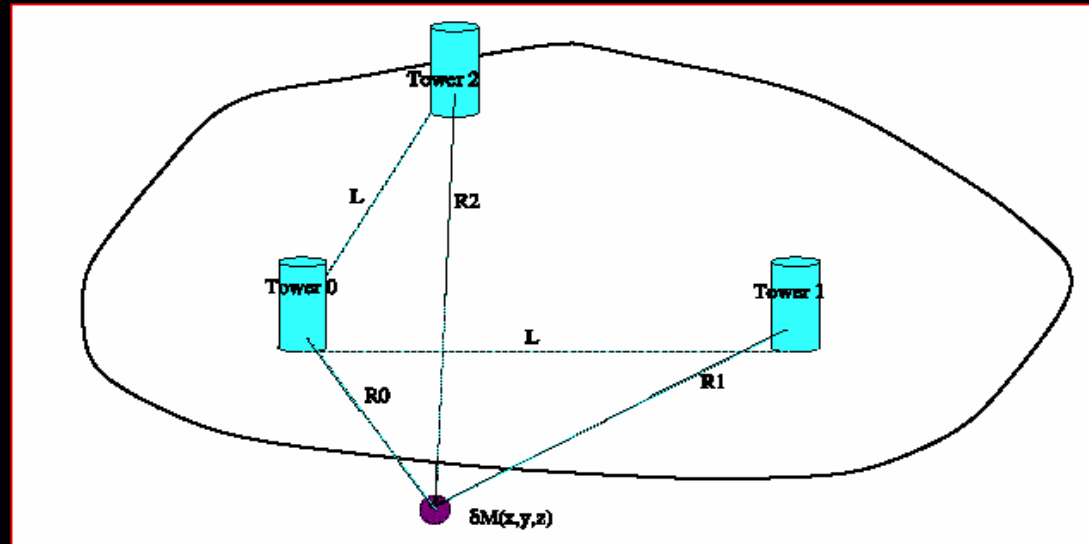
- Motivations
 - Seismic GGN
 - Atmospheric GGN
- Going underground
 - Underground GGN estimates
 - GGN reduction inside a cavity
- Other (but related) options
 - Monitoring and subtraction
 - Reference masses
- Conclusions



What is Gravity Gradient Noise

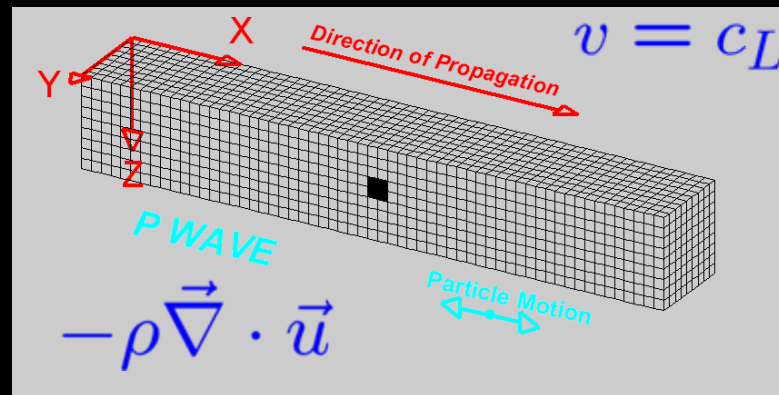
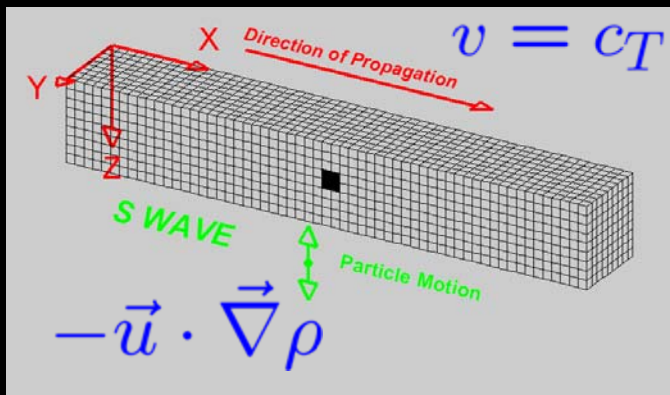


Mass density fluctuations couple directly to the test masses:



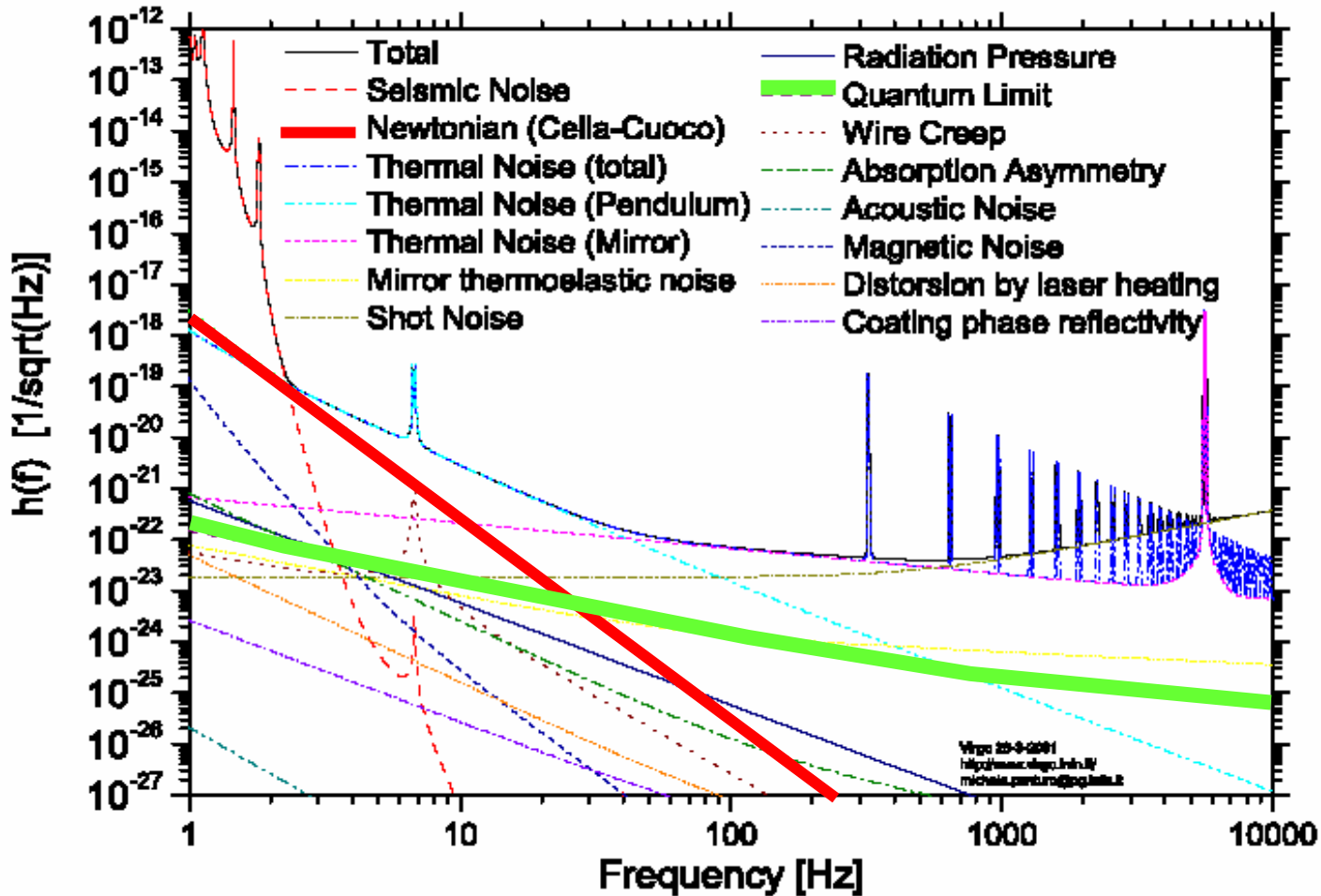
Example:
Elastic material

$$\delta\rho(x, t) = -\vec{\nabla} \cdot (\rho(x, t)\vec{u}(x, t))$$



$$\nu = \left(\frac{c_T}{c_L} \right)^2$$

Seismic GGN



Estimate uses transfer function between seism and GGN

$$seism \sim f^{-2}$$

$$TF \sim f^{-2}$$

$$GGN \sim f^{-4}$$

Atmospheric GGN



Rayleigh Bernard Scenarios (G.C., E. Cuoco, P. Tomassini)

$$\begin{aligned} \partial_t \vec{u} + (\vec{u} \cdot \vec{\nabla}) \vec{u} &= -\frac{1}{\rho} \vec{\nabla} p + \nu \nabla^2 \vec{u} + \alpha \vec{g} \theta \\ \partial_t \theta + (\vec{u} \cdot \vec{\nabla}) \theta &= \chi \nabla^2 \theta \quad \vec{\nabla} \cdot \vec{u} = 0 \end{aligned}$$

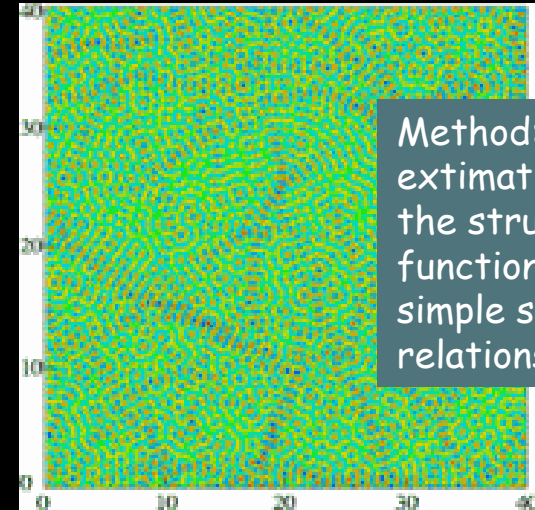
Different possibilities accordingly with the intensity of thermal gradient.

Thermal bubbles

- ✘ Nucleation phase (slow)
- ✘ Ascension phase

$$\begin{aligned} T_b(z) &= T_b(0) + \gamma_{ad} z \\ \partial_{tt} z &= g \left(\frac{T_b(z) - T(z)}{T(z)} \right) - 6\pi\nu r z \end{aligned}$$

Well developed turbulence

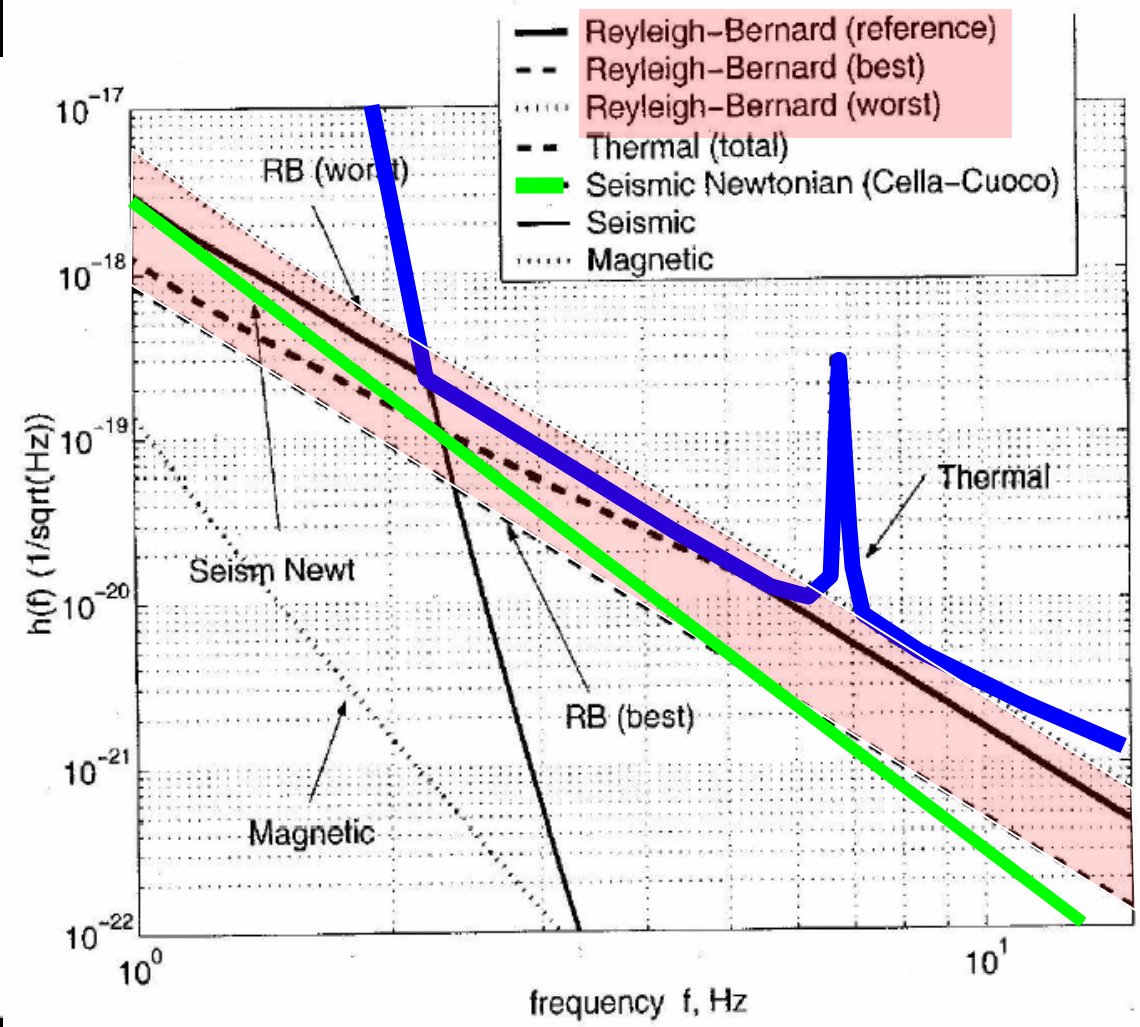
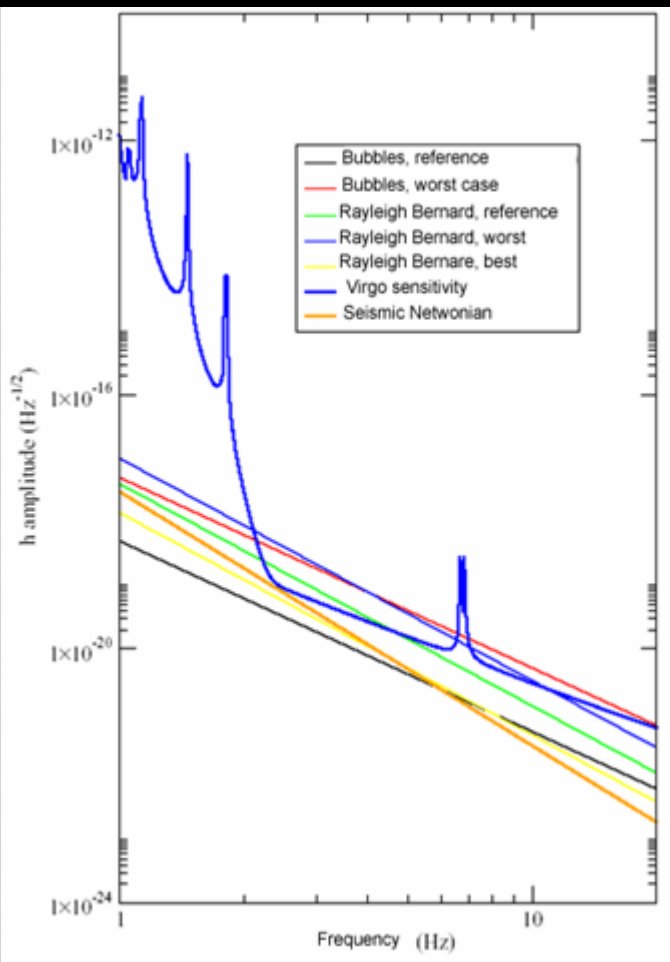


Method: estimation of the structure functions using simple scaling relations.

Other

- ✘ Effect of acoustic waves (Saulson)
Negligible
- ✘ Airborne objects, sonic booms, advection, ... (Creighton)
- ✘ Turbulent generation of acoustic waves (Lighthill process). (C. Cafaro, G. C.)
Negligible

Atmospheric GGN



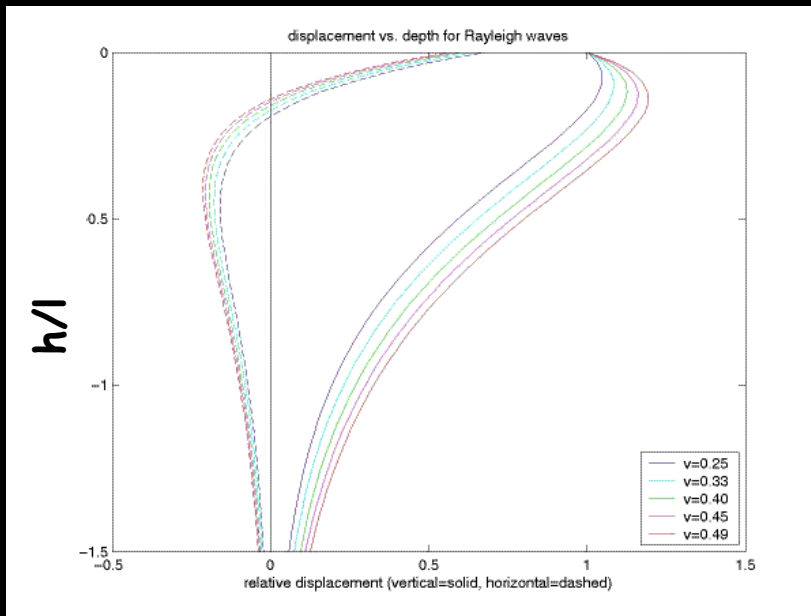
What we can expect by going underground



Going underground: seismic GGN reduction



A simple fact: surface waves die off exponentially with the depth



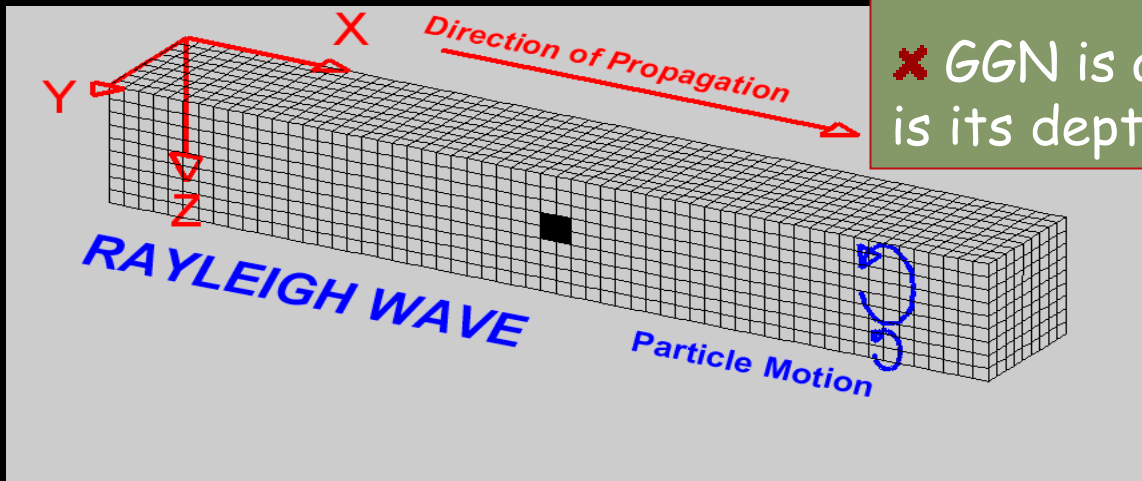
✘ Surface waves are probably the most important excitations for GGN

✘ Surface movement dominate the bulk compression effect

✘ Most efficient mechanism to transport energy from "far" sources

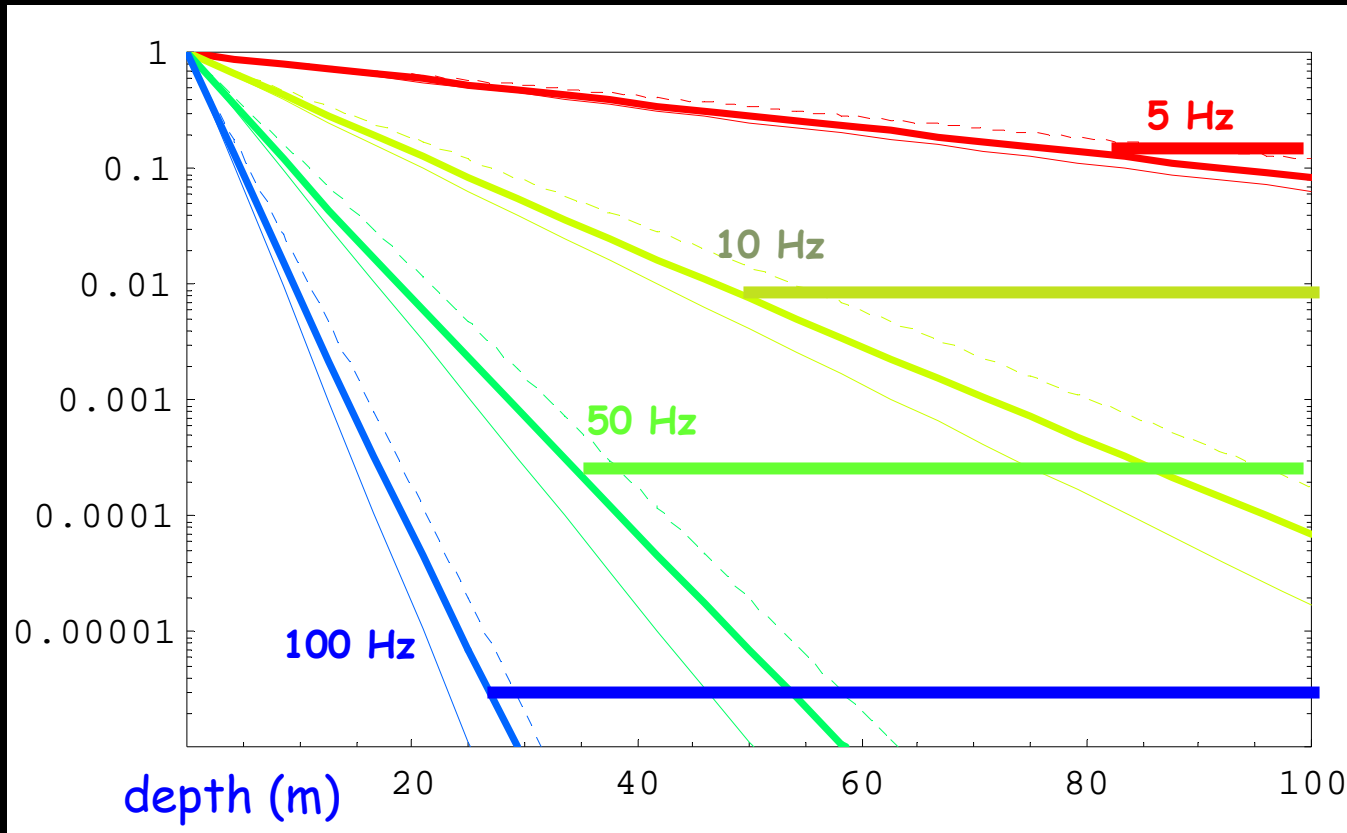
✘ Significant coupling with "local" sources (human activity)

✘ GGN is a "long range" effect: what is its depth dependence?



"Smearing effect"

GGN vs. depth



Relative reduction of GGN with depth:

Bulk

Surface

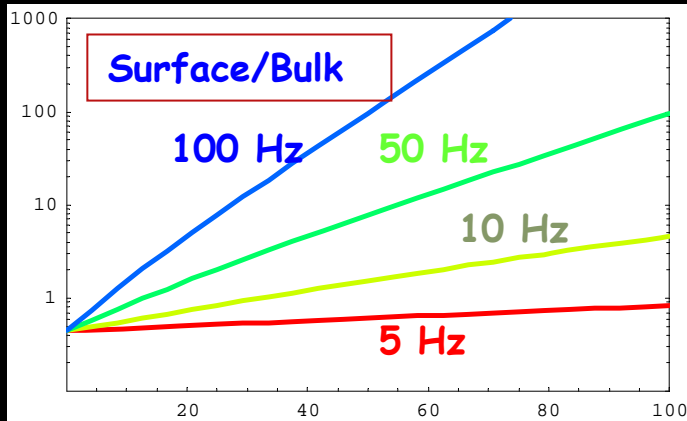
Total

$$bulk = g_1(f, \nu) e^{-h/\ell} + g_2(f, \nu) e^{-h/\ell_L}$$

$$surf = g_3(f, \nu) e^{-h/\ell}$$

$$\ell = \frac{c_T}{2\pi f} \sqrt{x}$$

$$\ell_L = \frac{c_T}{2\pi f} \sqrt{\frac{x}{1-x\nu}}$$



All this is pretty good, but

✗ Volume waves contributions will not share this fast decay

✗ Surface fluctuations in the depth?

A rough model for an underground cavity



Spherical cavity in a homogeneous elastic medium:

- ✘ Elasticity eq. $-\rho\omega^2\vec{u} = \mu\nabla^2\vec{u} + (\lambda + \mu)\vec{\nabla}(\vec{\nabla} \cdot \vec{u})$
- ✘ Free boundaries $\sigma \cdot \hat{n} = 0$
- ✘ Mode Classification accordingly with rotation symmetry:

$$\vec{u}_{l,m} = \underbrace{\vec{\nabla}\xi_{l,m}}_{\text{Spheroidal longitudinal}} + \underbrace{\vec{\nabla} \times (\vec{x} \times \vec{\nabla})\eta_{l,m}}_{\text{Spheroidal transverse}} + \underbrace{(\vec{x} \times \vec{\nabla})\tau_{l,m}}_{\text{Toroidal}}$$

For each ω, l, m :

- ✘ 2 spheroidal modes (mixed transverse & longitudinal)
- ✘ 1 toroidal mode (transverse only)
- ✘ Incoming wave scattered to an outgoing one

$$\xi_{l,m} = Y_{lm}(\theta, \phi)\alpha_{l,m}^{(+)}h_l^{(+)}(qr) + \alpha_{l,m}^{(-)}h_l^{(-)}(qr)$$

$$\eta_{l,m} = Y_{lm}(\theta, \phi)\gamma_{l,m}^{(+)}h_l^{(+)}(kr) + \gamma_{l,m}^{(-)}h_l^{(-)}(kr)$$

$$\tau_{l,m} = Y_{lm}(\theta, \phi)\beta_{l,m}^{(+)}h_l^{(+)}(kr) + \beta_{l,m}^{(-)}h_l^{(-)}(kr)$$

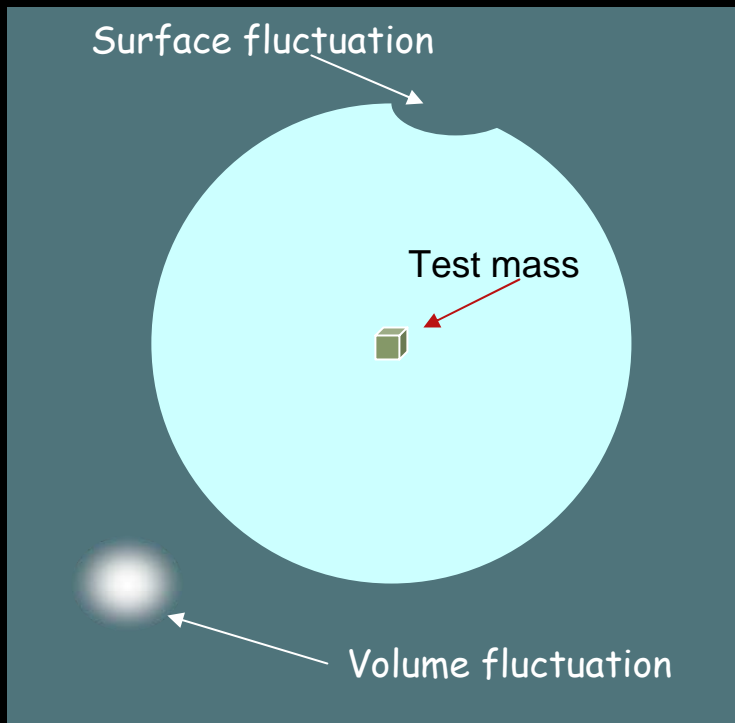
$$k^2 = \rho\omega^2/\mu \quad q^2 = \rho\omega^2/(\lambda + 2\mu)$$

What is the contribution of each mode to GGN?

A rough model for an underground cavity



$$\vec{u}_{l,m} = \vec{\nabla}\xi_{l,m} + \vec{\nabla} \times (\vec{x} \times \vec{\nabla})\eta_{l,m} + (\vec{x} \times \vec{\nabla})\tau_{l,m}$$



Bulk contribution to GGN:

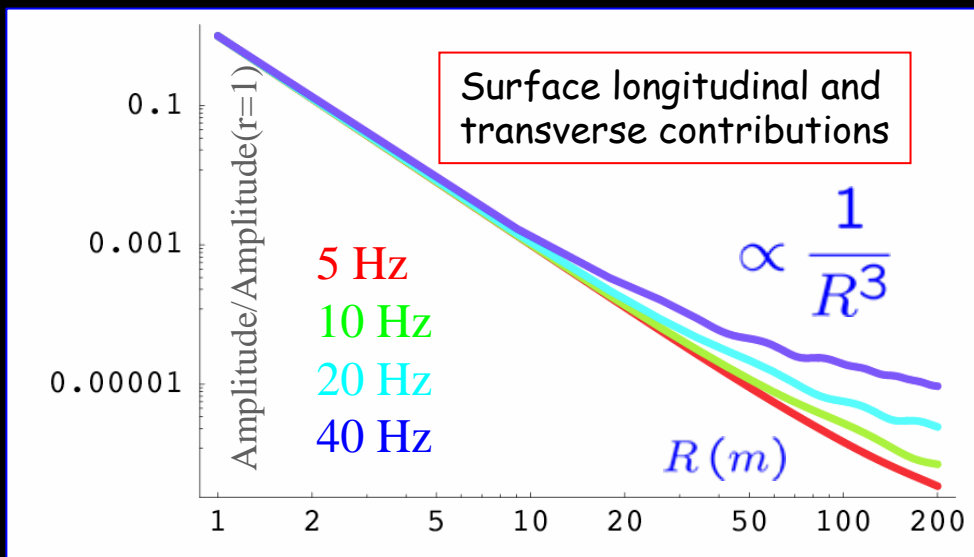
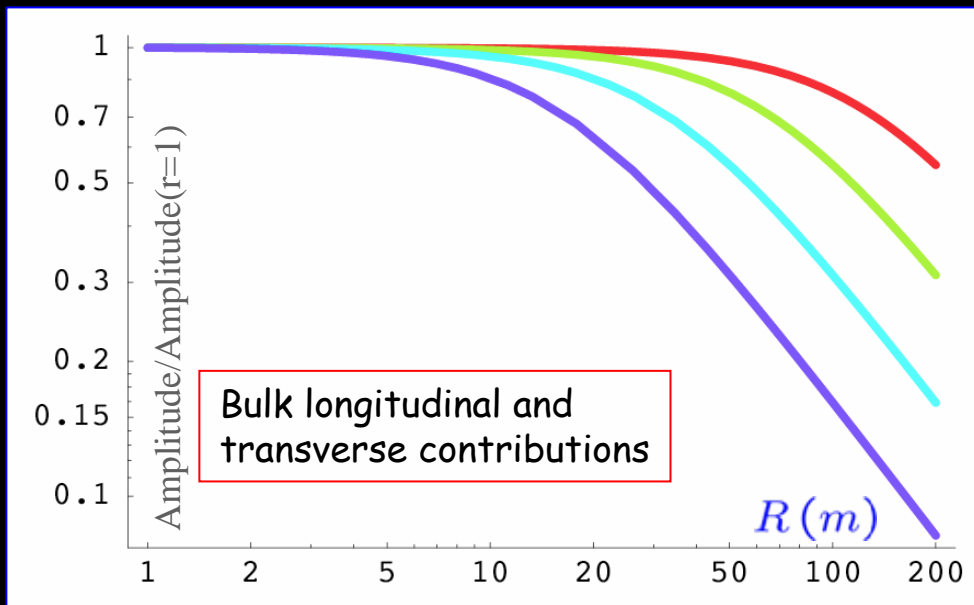
$$\vec{a}_{l,m} = -G\rho \int \vec{\nabla} \cdot \vec{u}_{l,m} \frac{\vec{r}}{r^3} dV$$

Surface contribution to GGN:

$$\vec{a}_{l,m} = -G\rho \int \vec{R} \cdot \vec{u}_{l,m} \frac{\vec{R}}{R^2} d\Omega$$

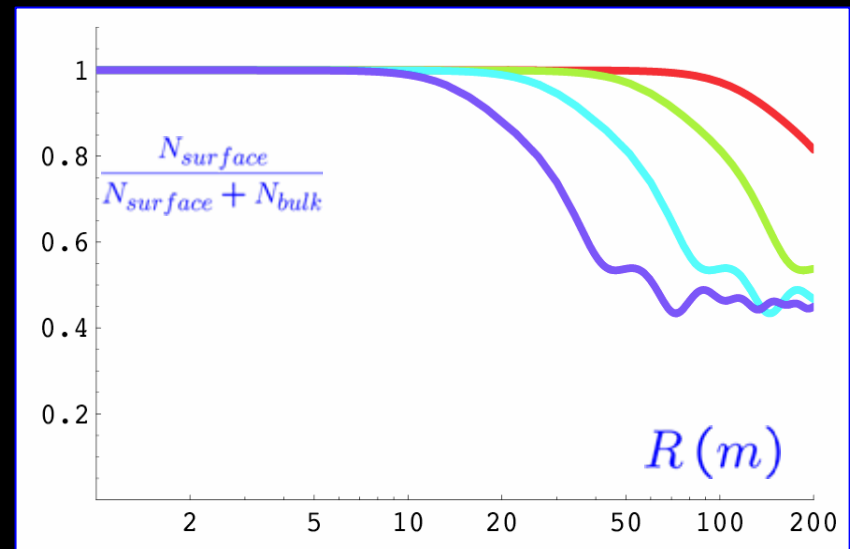
- ✗ Only "dipole" contribution $\sim Y_{1m}(\theta, \phi)$ to bulk GGN (cavity displacements)
- ✗ Both transverse & longitudinal contributions to surface GGN
- ✗ Toroidal modes: transverse, no surface motion, no Newtonian

R dependence of GGN inside the cavity



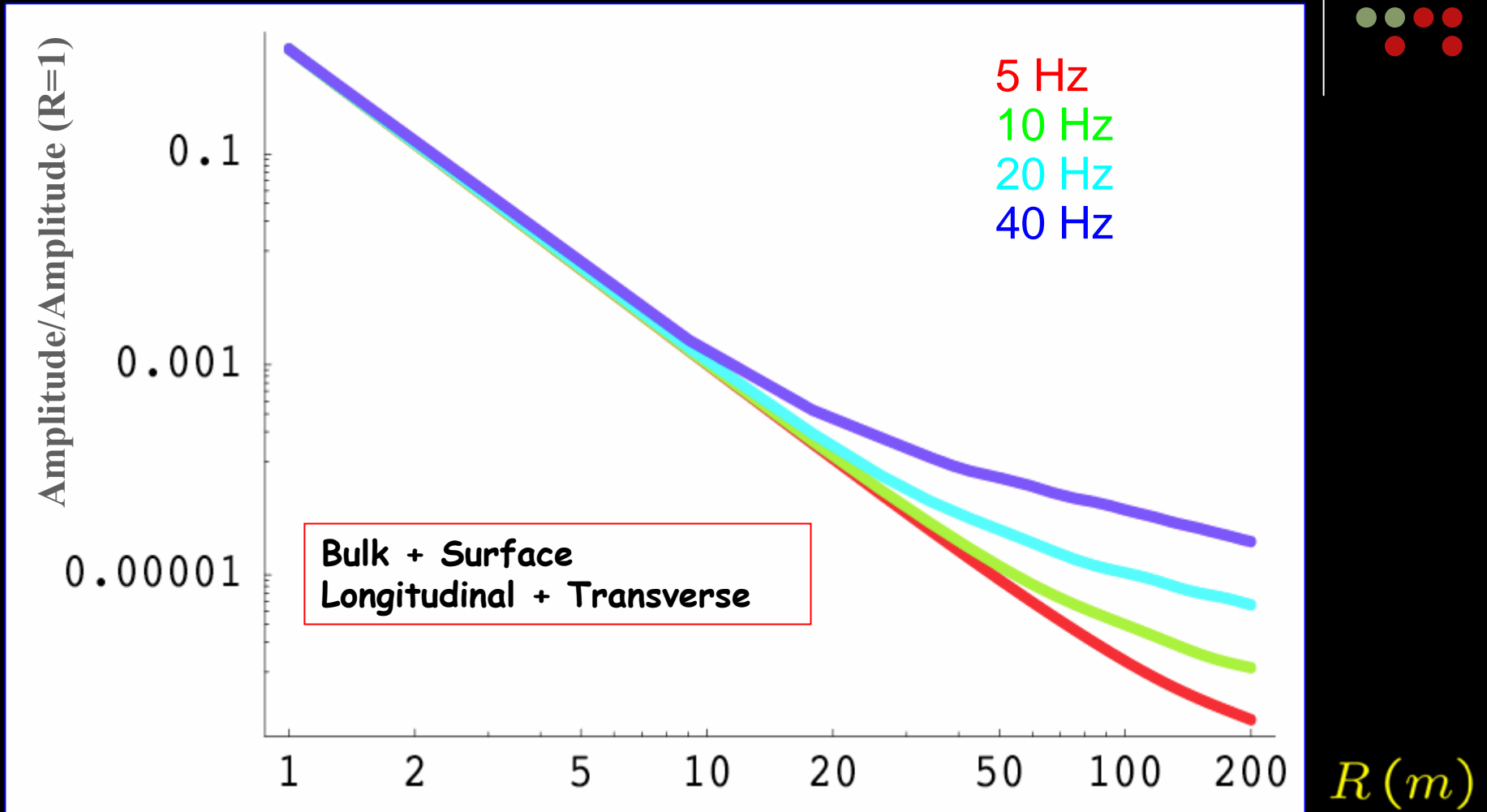
Method:

- ✗ Normalization: fixed incoming flux from infinity of elastic energy
- ✗ Statistical sum over modes
- ✗ Assumed absence of correlations between modes (weak dependence)



Surface contributions always dominant in the relevant frequency range

R dependence: final result



Good reduction with a reasonable cavity's size.

Transfer function



There is a relation between GGN in the cavity and seismic motion measured on the surface?

$$\hat{n} \cdot \vec{u}_{l,m} = \frac{\partial}{\partial r} \xi_{l,m} + \frac{l(l+1)}{R} \eta_{l,m} \quad \text{Motion normal to the surface (as an example)}$$

$$\langle |\hat{n} \cdot \vec{u}|^2 \rangle \sim \sum_{\ell} g_{\ell}^2 \frac{2\ell+1}{4\pi}$$

Symmetries are not constraining enough.....

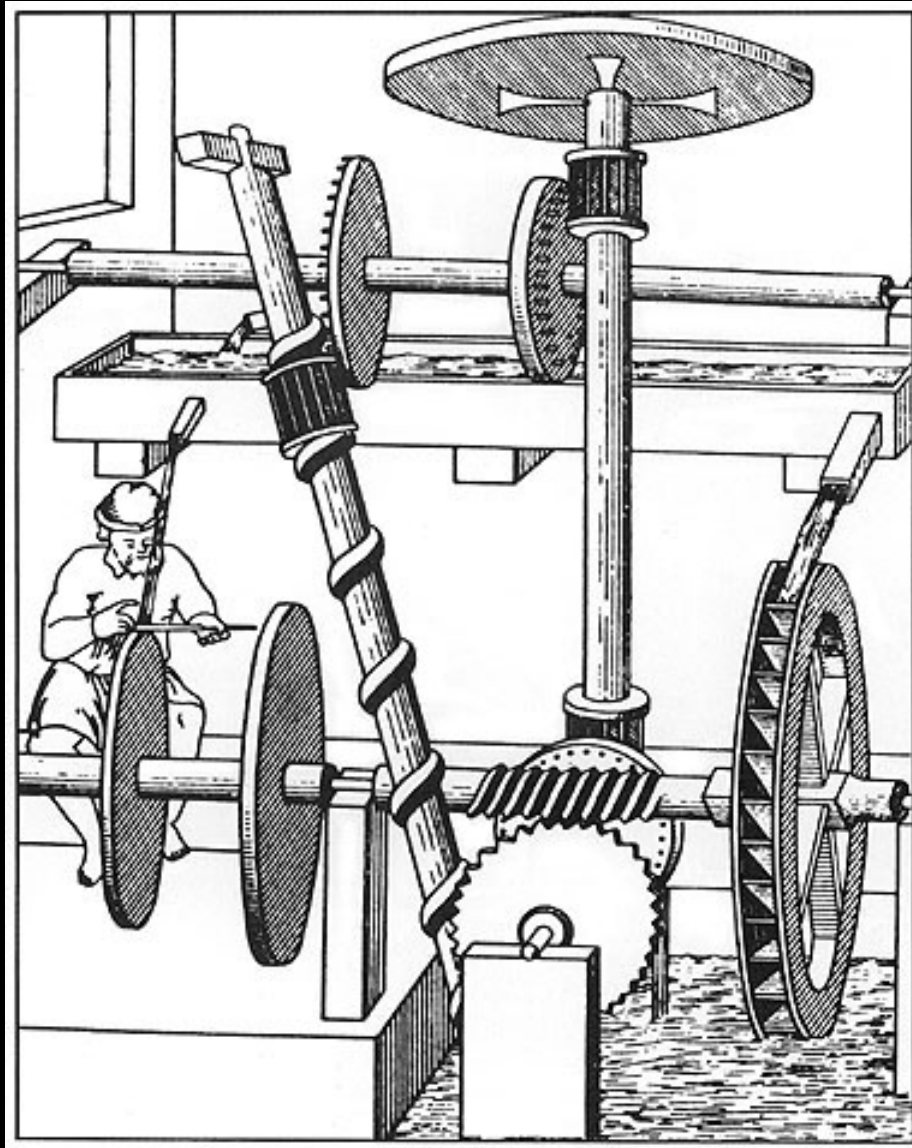
- Seismic motion get contributions from all g_{ℓ}
- GGN is controlled by $\ell = 1$ modes only

In other words: measuring the dipole mode is a difficult issue, without additional informations about the importance of each g_{ℓ} .

In principle: measure the correlations

$$\langle (\hat{n} \cdot \vec{u}) (\hat{n} \cdot \vec{u}') \rangle \sim \sum_{\ell} g_{\ell}^2 \frac{2\ell+1}{4\pi} P_{\ell}(\cos \gamma)$$

Other (related) options

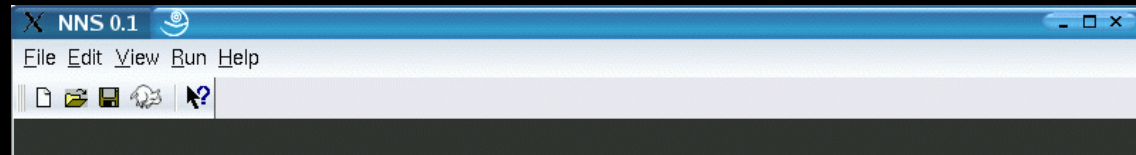
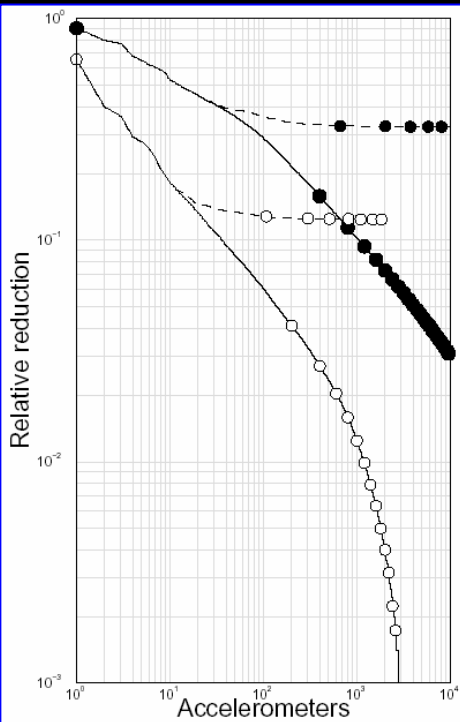


GGN subtraction

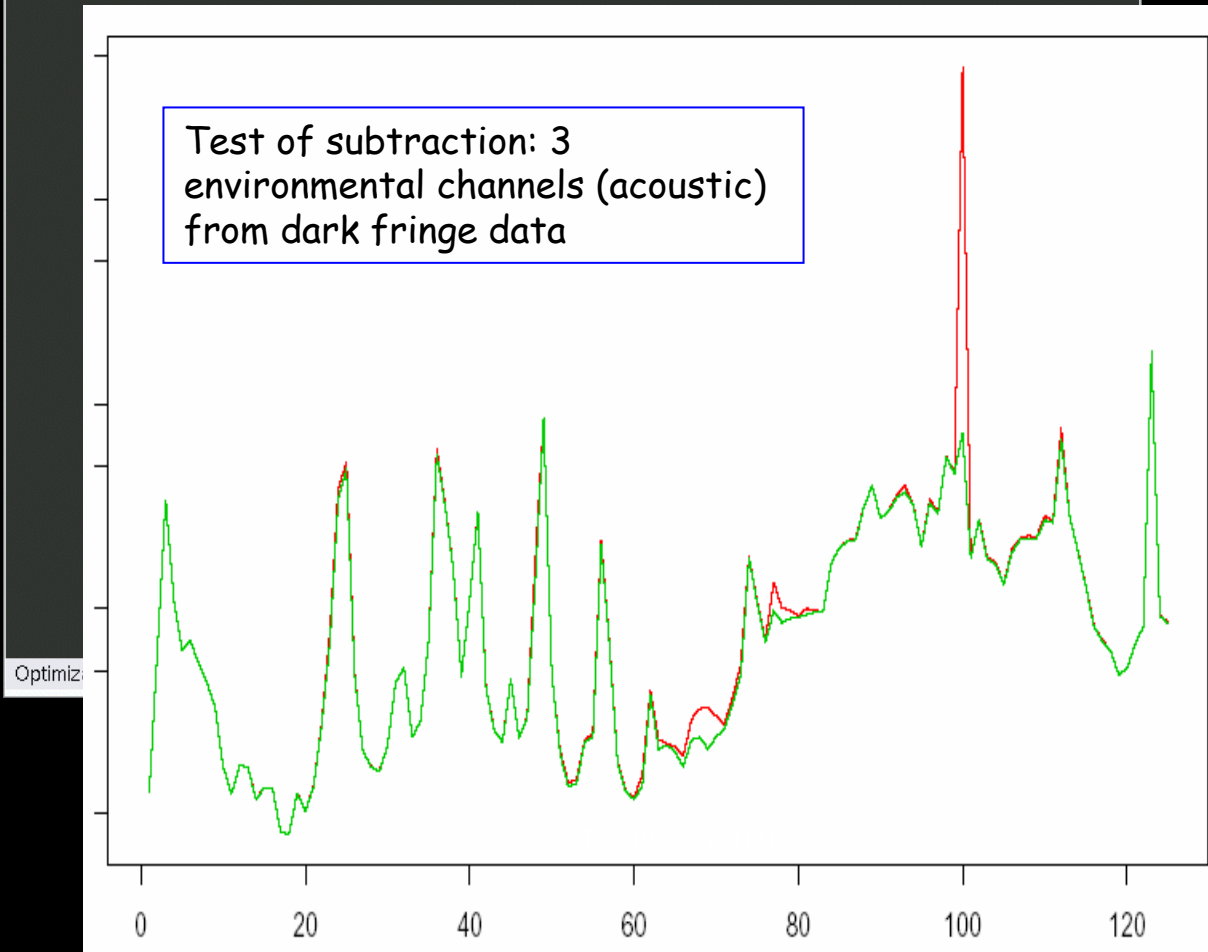


Example: a set of 10 accelerometers with random initial position....

...and after the optimization



Test of subtraction: 3 environmental channels (acoustic) from dark fringe data



Subtraction in the cavity?



We can apply the subtraction method to seismic measurements inside the cavity.

$$X_{sub}(\omega) = X(\omega) - \sum_{i,j} \langle X(\omega) \hat{n}_i \cdot \vec{u}(x_i, \omega)^* \rangle [C^{-1}(\omega)]_{ij} \hat{n}_j \cdot \vec{u}(x_j, \omega)$$

Subtracted signal

$$\eta = \frac{\sum_{i,j} \langle X(\omega)^* \hat{n}_i \cdot \vec{u}(x_i, \omega) \rangle [C^{-1}(\omega)]_{ij} \langle \hat{n}_i \cdot \vec{u}(x_i, \omega)^* X(\omega) \rangle}{\langle X(\omega)^* X(\omega) \rangle}$$

Subtraction efficiency

- The method can be applied
- We can't anticipate its efficiency
- Performances and number of sensors will depend on the number of relevant modes

Measuring horizontal GGN

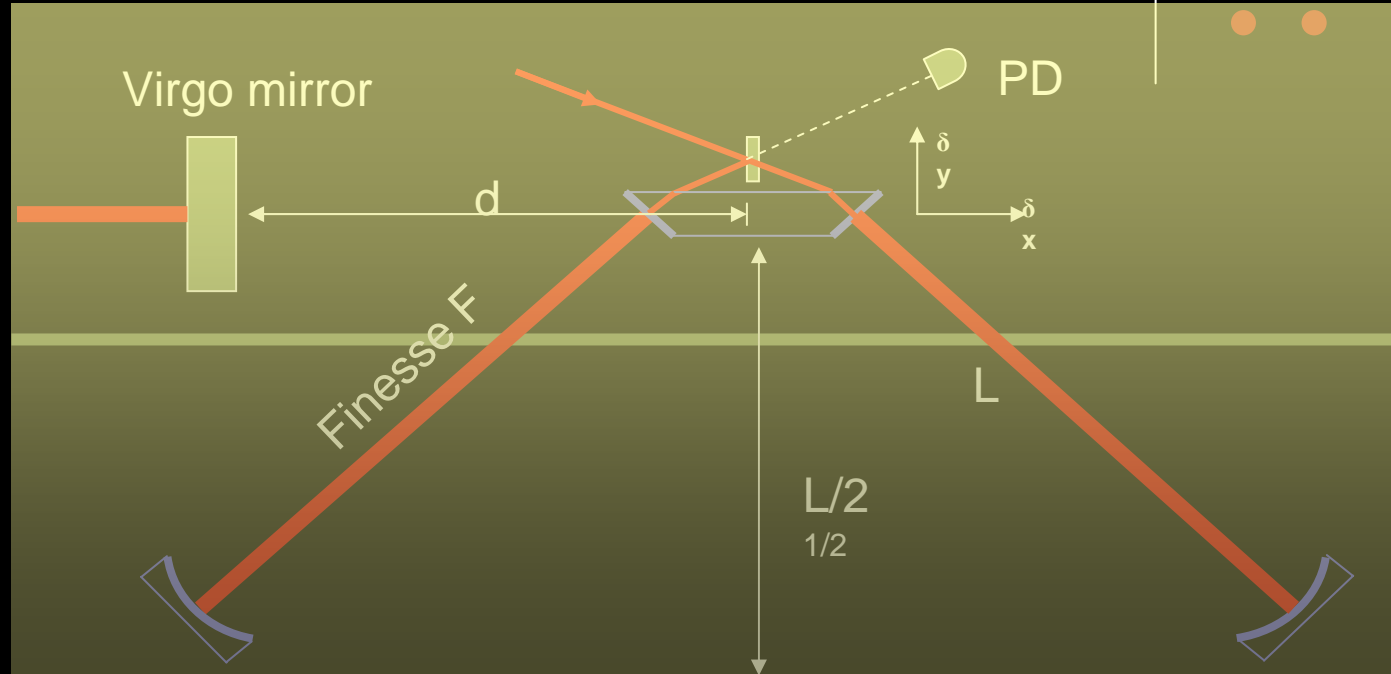


Preliminary idea:

1. Measure the GGN using as a reference "quiet" masses.
2. Subtract

Features:

1. Decoupled from vertical GGN
2. L large compared with
3. L small compared with interferometer arms



Problems:

1. "Quiet" masses must be dominated by GGN
2. "Quiet masses coupled to vertical seism (more refined schemes can cure this problem)

Conclusions:

- ✗ The underground option seems promising
 - ✗ Seismic surface waves contributions to GGN exponentially damped
 - ✗ Atmospheric contributions should be damped also exponentially
- ✗ A cavity can be used to further reduce GGN

Problems:

- ✗ Localized seismic waves on the gallery
 - ✗ Small masses involved
 - ✗ Monitorable
- ✗ Acoustic (pressure waves) resonances
 - ✗ Could be reduced
 - ✗ Monitorable
- ✗ Volume seismic waves

Measurements in realistic scenarios are mandatory!

Thank you for your attention!

